whole medium (Fig. 1, G and H); thus, from two spirals only one is left. (iv) The spirals drift toward higher light intensity, as predicted in (13), which can be seen from a comparison of the tip position at different times; the drift is screw-like, as it is accompanied by the rotating motion of the spiral tip. The mean velocity of the drift has one component in the direction of the light gradient and another one perpendicular to it. The mean drift velocity, which can be estimated by comparing Fig. 1G with Fig. 1H, is 1.7×10^{-2} cm/min. (v) Strong deviations from the Archimedean spiral occur, the wavelength becoming larger in the regions with lower light intensity.

Figure 2 shows the automaton simulations corresponding to the experiments of Fig. 1. In agreement with these experimental results, the simulations yield instability in which the symmetry of the spiral pair is lost; the wavelength is enhanced; a spiral pair is transformed into a single spiral; there is spiral drift; and the system deviates from the Archimedean shape.

In order to quantify the automaton results, it is necessary that the iteration step and the length of a cell be expressed in physical time and space units. This can be done by comparing the period T and the wavelength λ of a spiral in homogeneous light (Figs. 1A and 2A). Experiments yield T = 1.5 min and $\lambda =$ 0.23 cm. Simulations render T = 7 steps and $\lambda = 31$ cells. Thus, one iteration step corresponds to 0.21 min and one cell automaton is 7.4×10^{-3} cm long. The mean drift velocity, estimated by comparing Fig. 2G with Fig. 2H, is 0.53 cells per step, which in physical units is 1.9×10^{-2} cm/min, in good agreement with the experimental value of 1.7×10^{-2} cm/min obtained above.

An experiment with a radial gradient is shown in the upper pictures of Fig. 3; the corresponding automaton simulations are displayed below. Both in experiments and in simulations we observed that spirals straighten up, leading to uncurled wave ends. Also, we observed partial breaking of wave fronts into pieces. At the end of the experiment the mean drift velocity of the spiral tip is 5.9 $\times~10^{-2}$ cm/min, 3.5 times the velocity for the axial light gradient (Fig. 1). This is in good agreement with simulations. In fact, comparing the last two figures in the lower part of Fig. 3, we obtain a velocity of 2.1 cells per step, four times the value obtained for the axial gradient. This velocity enhancement is attributable to the fact that the spiral tip, while it is driven out of the medium in a screw-like fashion, encounters more often a driving light gradient when this gradient is radial than when the gradient points solely in one spatial direction.

The results presented here are at least qualitatively related to wave dynamics in heart muscle, as excitable media share some

common fundamental properties (1, 3). Measurements in vivo (2) and in isolated heart muscle (4) indicate that spiral formation is the cause of arrhythmic diseases. Moreover, simulations of heart muscle with partial differential equations (17) have shown that a gradient of refractoriness causes spiral drift. It has been suggested (18) that the traditional treatment, which consists of depolarizing shocks in the kilovolt range, may be replaced by implantable defibrillators operating at much lower potentials. In addition, the ability to control spiral drift in excitable media may find application in parallel processing with photochemical computational devices (9-11, 19).

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Kinetics of the OH Reaction with Methyl Chloroform and Its Atmospheric Implications

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The rate coefficients for the reaction of hydroxyl (OH) radicals with methyl chloroform (CH_3CCI_3) were measured between 243 and 379 kelvin with the pulsed photolysis–laser-induced fluorescence method. The measured rate coefficients at 298 and 277 kelvin were ~20 and ~15%, respectively, lower than earlier values. These results will increase the tropospheric OH concentrations derived from the CH_3CCI_3 budget analysis by ~15%. The predicted atmospheric lifetimes of species whose main loss process is the reaction with OH in the troposphere will be lowered by 15% with consequent changes in their budgets, global warming potentials, and ozone depletion potentials.

Methyl chloroform (CH_3CCl_3) is exclusively man-made. Its rate of release into the atmosphere can be accurately estimated from industry production figures (1), and its atmospheric loss is due primarily to reaction with OH radicals. [Recently, oceanic consumption has been identified as an additional small loss process accounting for about 7% of the total CH_3CCl_3 removal (2).] Therefore, measurements of atmospheric CH₃CCl₃ concentrations have been used to deduce globally averaged tropospheric OH concentrations (3), which in turn are used to calculate atmospheric life-

times of species such as methane (CH₄) and haloethanes, which are used as chlorofluorocarbon (CFC) substitutes. The OH concentrations calculated from the CH₃CCl₃ budget analysis critically depend on the rate coefficient, k_1 , for the reaction

$$OH + CH_3CCl_3 \rightarrow H_2O + CH_2CCl_3 \quad (1)$$

The generally accepted value of k_1 (4, 5) is based on the results of Kurylo *et al.* (6) and Jeong and Kaufman (7), which are in good agreement with each other between 278 and 363 K. Investigations carried out earlier than these two studies reported significantly

SCIENCE • VOL. 257 • 10 JULY 1992



Fig. 1. Plot of $\ln k_1$ versus 1/T measured in this study (open circles), by Kurylo *et al.* (6) (triangles), and by Jeong and Kaufman (7) (diamonds). The solid line is a fit of our data to an Arrhenius expression. The dashed line represents the recommendations of the National Aeronautics and Space Administration (4) and the International Union of Pure and Applied Chemistry (5) that are currently used in atmospheric calculations.

higher values, which were erroneous for reasons discussed by Kurylo *et al.* and Jeong and Kaufman.

The direct determinations of k_1 , that is, without reference to another rate coefficient, have been carried out using the flow tube or pulsed photolysis method. Because in these methods the loss of OH reactant is measured in an excess of CH₃CCl₃, there are two potential sources of error. First, stabilizers and contaminants in the CH₃CCl₃ sample can also react with the OH radicals. Stabilizers, such as 1,4-dioxane, are added to CH₃CCl₃ samples to prevent its decomposition and they react rapidly with OH. The common contaminant CH_2CCl_2 is a byproduct of CH₃CCl₃ production and is also produced by its decomposition; hence, it is usually present in CH₃CCl₃ samples. The rate coefficient for the reaction of OH with CH_2CCl_2 is >10⁴ times k_1 . Second, significant concentrations of free radicals may be

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produced by the reaction of CH_3CCl_3 with OH. These free radicals can enhance the observed OH removal rates; if so, the measured rate coefficients will be higher than the true values (8). In the pulsed photolysis experiments, there may be an additional complication due to the photolysis of CH_3CCl_3 and the reactions of photolytic products with OH.

We measured k_1 using the pulsed photolysis-laser-induced fluorescence method, taking great care to identify and minimize the problems mentioned above. The apparatus and experimental procedures are described elsewhere (9). We used various OH sources (10, 11): (i) 355-nm laser photolysis of HONO, (ii) 248-nm laser photolysis of H₂O₂ and HNO₃, and (iii) broad-band (Xe flash lamp, wavelengths >165 nm) photolysis of HNO₃ and H₂O. CH₃CCl₃ does not absorb at 355 nm. Its absorption cross section at 248 nm is $\sim 2 \times 10^{-22}$ cm². Therefore, CH₃CCl₃ photolysis was not a problem in these studies. The third source was used to investigate if substantial differences would be observed when CH₃CCl₃ was also photolyzed.

We monitored the relative OH concentrations, by using pulsed laser-induced fluorescence, as a function of time between the photolysis and probe laser pulses. The concentration of CH₃CCl₃ was typically $>10^5$ times that of OH. The concentration of CH₃CCl₃ flowing through the reactor was measured by absorption at 214 nm (Zn lamp) in cells located ahead of and behind the reactor. Thus, the required CH₃CCl₃ concentration was directly measured. Any significant loss (>2%) of CH₃CCl₃ to the walls or to heterogeneous decomposition would have been detected. It is important to know this concentration because the uncertainty in k_1 is directly proportional to the uncertainty in the CH₃CCl₃ concentration. We measured the absorption cross section of CH₃CCl₃ at 214 nm, which is needed to determine its concentration, to be $(1.33 \pm 0.05) \times 10^{-19} \text{ cm}^2$. The uncertainty represents a combination of random and systematic errors at the 95% confidence level. The major systematic errors in this measurement are in the pressure measurements, the adsorption of CH₂CCl₂ on the absorption cell windows, and the presence of absorbing impurities. We measured the cross section using two absorption cells (100 and 25 cm long); the cell was either filled with CH₃CCl₃ or CH₃CCl₃ flowed through the cell at known pressures. In all cases, the Beer-Lambert law was strictly

Table 1. Summary of experimental conditions and the measured values of k_1 as a function of temperature. Quoted error bars are 2σ and do not include estimated systematic errors. Cell pressure, 25 to 75 torr; flow velocity through the interaction zone, 4 to 10 cm s⁻¹.

OH source	[OH] ₀ (10 ¹⁰ cm ⁻³)	[CH ₃ CCl ₃] range (10 ¹⁵ cm ⁻³)	[CH ₂ =CCl ₂] (ppmv)	k ₁ (corrected) (10 ⁻¹⁴ cm ³ molecule ⁻¹ s ⁻¹)
HONO	36	24–101	8	0.304 ± 0.019
H ₂ O ₂	8	25–83	12	0.267 ± 0.018
HONO	38	17–100	31 .	0.270 ± 0.019
H ₂ O ₂	6	23–90	7	0.349 ± 0.016
HONO	34	17–119	42	$\begin{array}{l} 0.355 \pm 0.026 \\ 0.346 \pm 0.030 \\ 0.410 \pm 0.013 \\ 0.405 \pm 0.017 \\ 0.377 \pm 0.018 \end{array}$
HONO	19	11–98	31	
H ₂ O ₂	11	15–107	12	
HONO	12–29	30–110	2	
H ₂ O ₂	4	22–104	33	
HONO	44	23–122	8	0.522 ± 0.030
H ₂ O ₂	5	14–93	19	0.520 ± 0.022
HONO	16	11–126	12	0.579 ± 0.022
HONO	23	14–114	42	0.530 ± 0.026
H ₂ O ₂	7	14–85	30	0.609 ± 0.030
H ₂ O	10	8–101	15	$\begin{array}{l} 0.898 \pm 0.064 \\ 0.899 \pm 0.100 \\ 0.910 \pm 0.030 \\ 0.859 \pm 0.027 \\ 0.851 \pm 0.023 \end{array}$
H ₂ O ₂	7	14–80	20	
H ₂ O2	4	16–92	40	
HONO	27	11–123	- 12	
HONO	23	13–94	20	
HONO	16	10–80	2	$\begin{array}{rrr} 1.49 & \pm \ 0.05 \\ 1.44 & \pm \ 0.10 \\ 1.65 & \pm \ 0.05 \end{array}$
H ₂ O	10	10–85	9	
H ₂ O ₂	7	12–78	20	
HONO	23	9–83	2	$\begin{array}{rrr} 2.07 & \pm 0.06 \\ 2.24 & \pm 0.11 \\ 2.08 & \pm 0.08 \end{array}$
H ₂ O	10	9–77	5	
HONO	31	10–97	20	
HNO ₃	20	5–59	6	$\begin{array}{rrrr} 2.67 & \pm & 0.05 \\ 2.71 & \pm & 0.05 \\ 3.05 & \pm & 0.11 \end{array}$
HONO	29	9–70	2	
H ₂ O	10	6–71	20	
	$\begin{array}{c} OH\\ \text{source} \\ \\ HONO\\ H_2O_2\\ HONO\\ H_2O_2\\ HONO\\ HONO\\ H_2O_2\\ HONO\\ H_2O_2\\ HONO\\ H_2O_2\\ HONO\\ H_2O_2\\ HONO\\ H_2O_2\\ HONO\\ H_2O_2\\ H_2O_2\\ HONO\\ H_2O_2\\ H_2O_2\\ HONO\\ H_2O_2\\ HONO\\ H_2O_2\\ HONO\\ HOO\\ HO$	$\begin{array}{c c} OH & [OH]_0 \\ (10^{10} \ cm^{-3}) \\ \hline \\ \text{Source} & (10^{10} \ cm^{-3}) \\ \hline \\ HONO & 36 \\ H_2O_2 & 8 \\ \hline \\ HONO & 38 \\ H_2O_2 & 6 \\ \hline \\ HONO & 34 \\ \hline \\ HONO & 19 \\ H_2O_2 & 11 \\ \hline \\ HONO & 19 \\ H_2O_2 & 11 \\ \hline \\ HONO & 12-29 \\ H_2O_2 & 4 \\ \hline \\ HONO & 12-29 \\ H_2O_2 & 4 \\ \hline \\ HONO & 12-29 \\ H_2O_2 & 7 \\ \hline \\ HONO & 16 \\ \hline \\ HONO & 23 \\ H_2O_2 & 7 \\ \hline \\ HONO & 23 \\ \hline \\ HONO & 23 \\ \hline \\ HONO & 23 \\ \hline \\ HONO & 16 \\ \hline \\ H_2O & 10 \\ \hline \\ H_2O_2 & 7 \\ \hline \\ HONO & 23 \\ \hline \\ HONO & 23 \\ \hline \\ HONO & 10 \\ \hline \\ H_2O & 10 \\ \hline \\ HONO & 23 \\ \hline \\ HONO & 29 \\ \hline \\ H_2O & 10 \\ \hline \\ HONO & 29 \\ \hline \\ H_2O & 10 \\ \hline \end{array}$	$\begin{array}{c c c c c c c c c c c c c c c c c c c $	$\begin{array}{c c c c c c c c c c c c c c c c c c c $

SCIENCE • VOL. 257 • 10 JULY 1992

Table 2. Comparison of results from recent investigations of k_1 and currently recommended values. All errors are those quoted by the authors.

k _{1(298,K)} (10 ^{−14} сm ³ s ^{−1})	A (10 ⁻¹² cm ³ s ⁻¹)	(<i>E/R</i> ± Δ <i>E/R</i>) (K)	7 range (K)	Technique*	Reference
$\frac{1.16 \pm 0.10^{\dagger}}{1.12 \pm 0.21^{\dagger}}$ $\frac{1.30 \pm 0.30}{1.09 \pm 0.35}$	5.41 ± 1.84 5.40 ± 1.8 5.4 ± 3.0	1831 ± 95 1810 ± 100 1803 ± 450	278–457 222–363 359–402 298	DF-RF FP-RF PR-UVKS Relative rate‡	(7) (6) (16) (16)
1.2 1.2 0.95 ± 0.08	5.0 5.1 1.75 ± 0.34	1800 ± 200 1800 ± 200 1550 ± 60	250–460 250–460 233–379	Rec. Rec. PP-LIF	(4) (5) This work

*DF-RF, discharge flow-resonance fluorescence; FP-RF, flash photolysis-resonance fluorescence; PR-UVKS: pulsed radiolysis-ultraviolet kinetic spectroscopy; PP-LIF, pulsed photolysis-laser-induced fluorescence; Rec., recommended. †These are scaled to 298 K using E/R = 1550 K. #Measured relative to the reaction of OH with CH₃Cl assuming a rate coefficient for that reaction of 1.36 × 10⁻¹⁴ cm³ molecule⁻¹ s⁻¹.

obeyed, and the results agreed to within 0.5%. The capacitance manometer used for pressure measurements is accurate to better than 1%. The levels of impurities in our samples were in the parts per million by volume (ppmv) range. Therefore, we conservatively estimate the uncertainty in the cross sections as ~3.5%. Our value is close to that reported earlier (12) [(1.23 \pm 0.02) \times 10⁻¹⁹ cm²]. We used the measured value to calculate the CH₃CCl₃ concentrations.

Nonstabilized CH₃CCl₃ obtained from Dow Chemicals was purified by passing He through it until >70% of the sample was removed. The main impurities in our sample, CH2CCl2, CHCl2CH3, and CHClCCl₂, were identified and quantified by gas chromatography (with flame ionization or mass spectrometric detection). The levels of CH₂CCl₂, the major impurity, varied between 2 and 42 ppmv in different samples and also with time in the same sample. For measurements below 298 K (where the interference from CH_2CCl_2 is more pronounced), we frequently checked for the levels of CH₂CCl₂. The rate coefficient for the OH + CH₂CCl, reaction $[(2.36 \pm 0.16) \times 10^{-12} \exp[(450^{\circ})]]$ \pm 20)/T] cm³ s⁻¹] was measured recently in this laboratory (13) under the pressure and gas composition conditions used here and was used to correct the measured values of k_1 . The corrections varied with the sample purity, increased with decreasing temperature (T), and were negligible above 298 K. Typically, the corrections were a few percent. In one case the correction was as high as 17%. The levels of CHCl₂CH₃ and CHClCCl₂ were always less than 10 ppmv and, on the basis of the rate coefficient for their reactions with OH (14), they contributed negligibly to the measured k_1 .

The values of k_1 measured between 233 and 379 K (Table 1 and Fig. 1) were independent of total pressure (25 to 75 torr), flow rate (4 to 10 cm s⁻¹), photolysis laser fluence (5 to 20 mJ cm⁻² at 248 nm and 15 to 40 mJ cm⁻² at 355 nm), and the OH sources. The data between 243 and 379 K are well represented by the expression: $(1.75 \pm 0.34) \times 10^{-12} \exp[-(1550 \pm$ 60)/T] cm³ s⁻¹. (The quoted uncertainties are at the 95% confidence level for each of the parameters and do not include estimated systematic errors.) The above Arrhenius parameters are not significantly different from those obtained by an analysis of our data between 298 and 379 K and thus confirm the negligible contributions of the impurity reactions to the measured value of k_1 . (Data obtained at 233 K were not included in this analysis because of the wider spread in the measured values and the larger difficulties associated with their measurements.) The parameters and uncertainties needed to calculate values of k_1 for atmospheric modeling are more appropriately presented (4, 5) as: $k_1(298 \text{ K}) = 9.5 \times 10^{-15} \text{ cm}^3 \text{ s}^{-1}$; $A = 1.75 \times 10^{-12} \text{ cm}^3 \text{ s}^{-1}$; E/R = 1550 K (where A is the Arrhenius A-factor, E is the activation energy, and R is the gas constant). The uncertainty, at the 95% confidence level, for the 298 K value is a factor f(298) = 1.08. At other temperatures the uncertainty factor f(T) for k_1 calculated from the above Arrhenius parameters is given by

$$f(T) = f(298 \text{ K}) \exp\left[(\Delta E/R)\left(\frac{1}{T} - \frac{1}{298}\right)\right]$$
(2)

where $\Delta E/R = 100$ K. The above error bounds account for the uncertainties in the ultraviolet (UV) absorption cross section at 214 nm, the difference between the value of k_1 calculated from the Arrhenius parameters and the value measured near 298 K (after correction for small differences in temperature), and the uncertainty in the temperature dependence of k_1 . The $k_1(298$ K) quoted above is the average of the value measured near 298 K and that calculated from the Arrhenius expression (which agree to within 4%).

Our values at 298 K (Table 2 and Fig. 1)

SCIENCE • VOL. 257 • 10 JULY 1992

and 277 K (Fig. 1) are 22 and 16%, respectively, lower than the recommendations (4, 5) that are based on results of Kurylo et al. (6) and Jeong and Kaufman (7). However, they do overlap with the recommended values within their quoted uncertainties. Our results are based on analyzed samples, are corrected for impurity contribution, and are based on directly measured CH₃CCl₃ concentrations. In addition, we kept the concentrations of OH and the CH₃CCl₃ photolysis products low enough to make secondary reactions negligible. We also avoided purification that involved heating CH₃CCl₃. We have observed the decomposition of CH₃CCl₃ to form CH₂CCl₂ at about 400 K in metal tubes.

Attempts were made to understand the reasons for the discrepancies between our results and those of Kurylo et al. (6), who used a technique similar to ours. We photolyzed HNO3 using a Xe flash lamp to produce OH, the same method as that used by Kurylo et al.; the OH concentration rose and then fell as a function of reaction time when CH₃CCl₃ was present. The values of k_1 obtained from the decay parts of the temporal profiles were considerably lower than those in the laser photolysis experiments. Secondary generation of OH interfered with measuring k_1 , and we were unable to reproduce the results of Kurylo et al. Recent experiments (15) have shown that HNO₃ photolysis at 193 nm produces very little OH but large amounts of O atoms. Therefore, vacuum UV photolysis of HNO3 is not a good source of OH for kinetic measurements. It is not clear if production of O atoms could have caused systematic errors in the study of Kurylo et al.

The measurements by Nelson *et al.* (16) cover only a small temperature range (359 to 402 K), and k_1 was not directly measured at or below room temperature. Their relative measurement at 298 K yields a lower value of k_1 than that calculated from their direct measurements.

The increasing discrepancy between our results and those of Kurylo et al., Jeong and Kaufman, and Nelson et al. with increasing temperature is opposite to what is expected if reactive impurities were present in CH₃CCl₃ samples used by these investigators. One possible cause for the disparity is the decomposition of CH₃CCl₃ at higher temperatures to produce the reactive CH₂CCl₂ in earlier studies. The decomposition of CH₃CCl₃ is quite rapid over Al surfaces and is the main reason for adding inhibitors to the commercial samples of CH₃CCl₃. Also, we observed decomposition of CH₃CCl₃ when it was heated in stainless steel tubes up to 400 K for a few minutes. (We found that CH₃CCl₃ samples stored in glass containers at 273 K were stable for at least 8 months as shown by the invariance of k_1 measured using the same sample but 8 months apart.) In some of the earlier studies, CH₃CCl₃ may have decomposed upon heating or on metal cell walls. We also have observed that the formation of CH₂CCl₂ is enhanced in the presence of water adsorbed on surfaces. Nelson *et al.* (16) used approximately 12 torr of H₂O to produce OH, and they could have inadvertently generated CH₂CCl₂.

Our results indicate that the previously accepted values of k_1 are too high above 240 K. The lower values of k_1 at tropospheric temperatures increase the OH concentrations estimated from the CH₃CCl₃ budget analysis. This increase, in turn, decreases the calculated tropospheric lifetimes of molecules whose primary loss is via reaction with OH. The magnitude of the change can be roughly gauged from the values of k_1 at 277 K, which is an approximate average of the temperature at which the OH-initiated degradation of a well-mixed tropospheric gas takes place. Our value of k_1 at 277 K is ~15% lower than the earlier recommendations. Examples of the implications of the new value of k_1 include (i) a decrease in the calculated atmospheric lifetime of CH₄ by 15% and a 15% increase in its flux to balance the larger atmospheric loss rate, and (ii) decreases in the calculated lifetimes of hydrofluorocarbons and hydrofluorochlorocarbons (HCFCs) and ozone depletion potentials of HCFCs, the proposed CFC substitutes, by 15%. In the case of CH_4 , the new value of k_1 will partially offset the balance in its budget that was calculated after our report on the OH + CH_4 reaction (8).

Note added in proof: DeMore (17) and Finlayson-Pitts et al. (18) have also recently measured k, to be substantially the same as that reported here

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1,1-dichloroethane cannot be larger than that for the reaction of OH with C_2H_6 . We have used the rate coefficient for the OH + C_2H_6 reaction (4) in placing the limit for the contribution of the OH + 1,1-dichloroethane reaction to the measured value of k_1 . The rate coefficient for the reaction of OH with CHCICCl₂ is from (4).

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Microaggregations of Oceanic Plankton Observed by Towed Video Microscopy

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Oceanic plankton have been hypothesized to occur in micropatches (<10 meters) that can have a large impact on marine ecosystem dynamics. Towed video microscopy was used to unobtrusively determine distributions of oceanic plankton over a continuum of scales from microns to hundreds of meters. Distinct, taxa-specific aggregations measuring less than 20 centimeters were found for copepods but not for nonmotile (cyanobacterial colonies) or asexual (doliolid phorozooids) forms, which suggests that these small patches are related to mating. Significant patchiness was also found on larger scales and was correlated among taxa, indicating physical control. These video observations provide new insights into basic plankton ecology by allowing quantitative assessment of individual plankton in their natural, undisturbed state.

The microscale (<10 m) environment of oceanic plankton is thought to be important in determining population and ecosystem dynamics (1, 2). Foraging models have shown that microscale patchiness in prey density can greatly enhance predator population growth (3, 4). In laboratory experiments, the abundance of prey required for maximal growth of predators is often well above that found in the field; the body size and the growth rate of field predators, however, are similar to those of laboratory animals grown in saturating food (1, 5). This discrepancy is hypothesized to be a result of field predators feeding in micropatches of prey that are too small to resolve with existing instrumentation (1, 4-6).

Spatial distributions in oceanic plankton on scales from microns to hundreds of meters were determined with the use of the newly developed video plankton recorder (VPR) (7). The VPR was equipped with conductivity, temperature, pressure, and flow sensors and was towed at 0.75 to 2.25 m/s near the surface (1 to 8 m) and outboard (7 m) of the side of the ship. Video recordings were scanned visually at 5 to 30 fields per second to detect the presence of organisms, and when they were found, taxa and time code were noted. Tow speed data from the flow sensor were used to convert the time code to distance along the tow path. Tows were made from the R.V. Oceanus on 29 and 30 August 1991 in continental shelf (70 m bottom depth; $40^{\circ}41'N$, $70^{\circ}33'W$) and slope ($39^{\circ}32'N$, $70^{\circ}00'W$) waters south of Woods Hole, Massachusetts. A tow from each area was analyzed for the presence of micropatchiness. Satellite infrared imagery together with VPR data on sea surface temperature and salinity indicated the slope station was positioned in the western edge of a warmcore Gulf Stream ring. Seas were calm, with the result that the wind-induced turbulent mixing rate was minimal, favoring micropatch formation (4).

The unobtrusive nature of the VPR (8) enabled observation and quantification of delicate forms that have typically been difficult or impossible to sample with traditional gear such as nets, pumps, and Niskin bottles. These forms included cyanobacterial colonies (Fig. 1A), marine snow, copepods carrying egg clutches, doliolids with buds (Fig. 1B), medusae, ctenophores, larvaceans (Fig. 1F), sarcodines, salp chains, and large, chainforming diatoms. Such forms can have a large impact on marine ecosystem dynamics, but little is known of their abundance and distribution. Large amounts of nitrogen-fixing cyanobacterial colonies, Trichodesmium sp. (Fig. 1A), were found along the transect in the warm-core ring (Fig. 2), which supports the view that they are potentially a major source of new nitrogen and organic carbon in open ocean areas (9). Likewise, gelatinous forms such as the tunicate Doliolum nationalis (Fig. 1B) were found in dense patches along the

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