Reports

Petrology of Venus: Further Deductions

Abstract. The carbon dioxide pressure of the atmosphere of Venus may be controlled by one or more of several decarbonation reactions. One of these is the formation of carbonate-rich liquids. The pressure-temperature equilibrium conditions of these reactions are in the range of those proposed for the surface of Venus

An analysis of the stability of mineralogic phases on the surface of Venus has recently been presented by Mueller (1). Such a priori considerations serve as a framework for formulating significant investigations to be carried out by planetary probes. The purpose of this report is to offer an extension of Mueller's ideas, to criticize several minor points in his paper, and to consider his suggestion that the high CO₂ content of the atmosphere is largely a result of the high temperature. The possible existence of pools of carbonate lava on the surface of Venus is discussed. The physicochemical condition assumed by Mueller is accepted as a basis for this discussion. The surface temperature is considered to be 425°C and the total pressure is considered to be equal to 10 atm with a relatively large amount of CO₂ and a small amount of H₂O.

Mueller indicates that the failure of the carbonation equilibria of MgSiO₃ cannot be invoked as an explanation for the supposed high CO₂ pressure in the lower atmosphere of Venus since formation of this phase by decarbonation at the temperatures assumed would result in CO₂ pressures of the order of hundreds of atmospheres. On the other hand, his statement that the CO₂ pressures over magnesite and dolomite are much greater than that supposed for the pressure of Venus cannot be interpreted rigidly since Weeks' (2) data indicate that a pressure of only 0.6 atm of CO₂ is required to maintain the stability of dolomite and 1.4 atm of CO₂ can maintain the stability of magnesite. Mueller's implicit suggestion, that the pressures required to maintain the stability of magnesium carbonates in the presence of silica is much greater than the assumed surface pressure of Venus, is, however, correct.

Bowen's decarbonation series (3), however, offers many carbonation equi-13 MARCH 1964

libria, some of which may occur at temperatures and pressures similar to those proposed for Venus, taking the part of "buffer" reactions. For example, extrapolation of the high-temperature experimental data (4) to 425°C indicates that the reactions

CaMgSi ₂ O	$_{3} + Mg_{2}SiO_{4}$	$+ 2CaCO_3 \longrightarrow$
Diopside	Forsterite	Calcite
		$3CaMgSiO_4 + 2 CO_2$
		Monticellite
and		

 $CaMgSi_2O_6 + 2CaCO_3 \longrightarrow$ Diopside Calcite $Ca_2MgSi_2O_7 + CO_2$ Akermanite

are, at this temperature, in equilibrium at approximately 10 atm. It may be that these reactions buffer the CO2 pressure of the atmosphere of Venus.

It must be remembered that the inflections of many pressure-temperature curves are strongest precisely in the region of 425°C. Thus minor differences in the values assumed for surface temperature or for the composition of vapor or solid phases may be reflected by considerable changes in the carbonation equilibria. Thus the physicochemical parameters become increasingly significant and the precise location of the pressure-temperature curves by experimentation becomes more important.

The existence of carbonate melts on the surface of Venus must also be considered. Wyllie and Tuttle (5) have found that such a melt can be obtained in the system CaCO₃-Ca(OH)₂ at temperatures as low as 630°C at a eutectic which is at a pressure of less than 27 atm. Similarly, Gittins and Tuttle (6) have discovered that the eutectic in the system Na₂CO₃-Ca(OH)₂ is at 595°C and that in the system Ca(OH) 2-Na2 CO₃-CaCO₃ at 1000 atm with 20 percent H₂O is at 490°C. These minimum melting temperatures will be lowered, perhaps considerably, by the addition of other components such as K₂CO₃ and even MgCO₃. On the other hand, the temperature of melting and the amount of liquid formed is strongly dependent on the amount of water in the vapor phase so that a mixture of CaCO₃ and vapor with equal parts of water and CO₂ will, at 1000 atm, begin melting only at about 1000°C. The effect of vapor composition on more complex systems is yet to be determined.

Recent descriptions of the active African volcano, Oldoinyo Lengai (7), state that it is ejecting carbonate lavas which are very rich in alkalies. Since the molten ejecta do not glow, a temperature less than 600°C is indicated. This occurrence also indicates that the pressure necessary to maintain a natural carbonate melt may be quite low.

It is not considered likely that the entire surface of Venus is covered by carbonate melt. The assumed temperature and low water content of the atmosphere would permit only the lowestmelting fraction of the rocks to form a liquid phase. In certain locations, where temperatures may be higher and the abundance of alkalies greater, the surface of Venus may have pools of carbonate liquid.

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References and Notes

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Xenon Tetroxide: Mass Spectrum

A new compound, xenon tetroxide, was formed by reaction of sodium perxenate (1) with concentrated sulfuric acid at room temperature. It was identified by mass analysis in a Bendix timeof-flight mass spectrometer with a modified source (2). The sample was prepared and introduced into the spectrometer by way of a pyrex system.

The spectrum of XeO4 is shown in the oscilloscope display of Fig. 1. The typical isotopic pattern of xenon is repeated every 16 mass units up to XeO4. The exact mass of the tetroxide is verified further by reference to the