full colleterial glands 24 to 31 days after the operation [Fig. 1 (3 and 4)]. When real or wax oöthecae were present in the uterus, oöcytes developed only after the nerve cord was severed.

In both *Blattella germanica* and *Pycnoscelus surinamensis,* inhibition of oöcyte development during pregnancy appears to be due to nervous stimuli resulting from pressure of the oötheca—on the uterus in *P. surinamensis* and on the oöthecal chamber in *B. germanica.* There was no indication that any substance released by the eggs in the oötheca acted through the brain to inhibit secretion of the *corpora allata,* as has been reported for *Leucophaea maderae* (*1*).

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4. Insects were kept at 24° to 25°C and at about 50 to 70 percent relative humidity.
5. The numbers following ± represent standard errors.

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Fault Zone along Northern Boundary of Western Snake River Plain, Idaho

Abstract. Gravity, seismic, and geologic studies indicate that at least 9000 ft of aggregate throw along a zone of northwest-trending, high-angle faults has displaced the western Snake River Plain downward relative to highlands on the north. At least 5000 ft of movement occurred between the early and middle Pliocene. Progressively diminishing movement since then amounts to 4000 ft.

Movement within a northwest-trending zone of high-angle faults, producing the abrupt escarpment that rises along the northern boundary of the western Snake River Plain in Idaho, was first recognized during reconnaissance surveys 60 years ago (*1*) but was disregarded in subsequent discussions that attributed the Snake River Plain to downwarping (*2*). Recently published gravity and seismic studies support the role of faulting as a cause of the escarpment, and current geologic mapping by the U.S. Geological Survey helps to date the times of movement. The fault zone aligns middle Pliocene and younger rocks of the Snake River Plain against a variety of older rocks in highlands to the north, including granite of the Idaho batholith north of Mayfield and silicic volcanics of early Pliocene age north of Mountain Home.

The pertinent gravity measurements are in a network of 800 locations that embraces southern Idaho and adjacent parts of neighboring states. In this network the greatest change in gravity is across the northern boundary of the Snake River Plain, amounting to a difference of 166 milligals between a high in the Snake River Plain 10 mi southwest of Mountain Home and the low in the Idaho batholith (*3*). This difference in gravity is opposite to that associated with high-angle faults elsewhere in western United States and indicates that rocks in the downthrown block of the Snake River Plain are heavy compared with the Idaho batholith. From analysis of a 50-milligal residual anomaly associated with the steep gravity gradient near Mountain Home, it is calculated that from 13,000 to 38,000 ft of rocks about as dense as Columbia River basalt have been dropped down against the Idaho batholith (*4*), the thickness of the displaced rocks depending on the density contrast assumed (0.3 and 0.1 g/cm^3, respectively).

The crustal break implied by the gravity measurements is possibly expressed by a line of earthquake epicenters that extends diagonally from Puget Sound, across the Columbia River Plateau, along the northern boundary of the western Snake River Plain, and thence across the plain to northern Utah (*5*). In Idaho, these earthquakes originate principally at average depths of 61 and 38 km (38 and 24 mi), the shallower earthquakes being near the base of the crust (*6*). The displacement calculated from the gravity measurements therefore ranges from one-tenth to one-third of the local crustal thickness.

Geologic evidence of intermittent crustal adjustment has been found near King Hill along the zone of faulting inferred from the geophysical observations, indicating that rocks of early Pliocene and younger age are dropped downward toward the Snake River Plain in progressively diminishing amounts. Thus, the bounding escarpment is formed by a large displacement of the silicic volcanics of early Pliocene age; basalt of middle Pliocene age that overlaps these faulted volcanics is offset a lesser amount; sedimentary rocks of late Pliocene age overlying the faulted basalt are somewhat offset and disarrayed by faulting; deposits of early Pleistocene age resting on the disturbed late Pliocene rocks are dropped down only a few hundred feet; and, finally, the deposits of middle Pleistocene and younger age are fractured, but not measurably offset, along the fault zone. The detailed paleontologic evidence by which this rock sequence has been dated will be described elsewhere, but the amounts of progressive crustal displacement can be summarized briefly.

The faulted silicic volcanics north of the fault zone are part of an eroded highland as much as 7500 ft above sea level, whereas drilling in the Snake River Plain 15 mi south of the fault zone reached a depth of 1200 ft below sea level without encountering these volcanics (*7*); this seemingly demonstrates a minimum displacement of about 9000 feet. The basalt of middle Pliocene age forms an upland 3800 ft above sea level north of the fault zone, but at King Hill, near the southern margin of the zone, drilling reached a depth of 1600 feet above sea level without penetrating this basalt (*8*); this suggests an offset of about 2200 feet. Displacement of the sedimentary rocks of upper Pliocene age is shown by exposures mainly along the northern margin of the fault zone, where one prominent fault offsets these rocks 600 ft, but tilting, no doubt caused by concealed faults within the fault zone, accounts for an equal amount of additional displacement, aggregating about 1200 ft. The deposits of early Pleistocene age along the fault zone are stepped down by a series of faults whose combined throw is about 600 ft. As the sum of all the measurable displacements in rocks younger than the silicic volcanics (4000 ft) accounts for less than half the minimum total displacement along the fault zone (9000 ft), most of the movement must have occurred before deposition of the middle Pliocene and younger rocks, but after eruption of the early Pliocene silicic volcanics (*9*).

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