ENSO variations change the basic state wind fields over the Atlantic at interannual time scales, altering environmental conditions that may affect hurricane formation. This relationship offers some accuracy in predicting the number of hurricanes that may occur in a given year. Variations of tropical winds and precipitation on intraseasonal time scales associated with the MJO (13, 19) are at least as large as the interannual variations associated with ENSO, and they may offer the possibility of predicting which periods during the hurricane season are most likely to produce hurricanes. The slow evolution of the MJO may be forecast up to 2 weeks or more into the future (27). An accurate MJO forecast, combined with knowledge of how the MJO affects Gulf of Mexico and Caribbean hurricanes, can be used to improve extended-range forecasts of tropical cyclone activity. The tendency of tropical cyclones over the Gulf of Mexico and Caribbean to cluster in time (28) may be explained at least in part by the interactions with the MJO described here.

References and Notes

- 1. Tropical storm, winds 34 to 64 knots (17 to 33 m s^{-1}); hurricane, winds greater than 64 knots.
- P. J. Hebert, J. D. Jarrell, M. Mayfield, NOAA Technical Memorandum No. NWS-NHC-31 (NOAA, Washington, DC, 1993).
- 3. P. Lester, *The Great Galveston Disaster* (Marrow, St. John, New Brunswick, Canada, 1900).
- 4. Category 5 on the Saffir-Simpson scale corresponds to winds greater than 135 knots (155 mph, 69 m s^{-1}); category 3 corresponds to winds of 96 to 113 knots (111 to 130 mph, 49 to 58 m s^{-1}).
- 5. R. H. Simpson and H. Riehl, *The Hurricane and Its Impact* (Louisiana State University Press, Baton Rouge, 1981).
- B. R. Jarvinen, C. J. Neumann, M. A. S. Davis, NOAA Technical Memorandum No. NWS-NHC-22 (NOAA, Washington, DC, 1984).
- 7. R. A. Madden and P. R. Julian, *Mon. Weather Rev.* **122**, 814 (1994).
- H. H. Hendon and M. L. Salby, J. Atmos. Sci. 51, 2225 (1994).
- 9. Anomalies are defined as deviations from the mean annual cycle.
- 10. E. Kalnay et al., Bull. Am. Meteorol. Soc. 77, 437 (1996). 11. Data not shown.
- 12. J. E. Kutzbach, J. Appl. Meteorol. 6, 791 (1967).
- 13. E. D. Maloney and D. L. Hartmann, unpublished data.
- E. D. Matchey and D. Z. Matchalli, unpublished data.
 EOFs were computed for the region: Equator-30°N, 80°W-130°W. The first EOF is shown in (13).
- The first PC has a zero-lag correlation of -0.6 with an MJO index in a previous study (29) that is dominated by western Pacific and Indian Ocean zonal wind variability.
- 16. Cyclonic circulations rotate in the same direction as the Earth, counterclockwise, as viewed from above in the Northern Hemisphere. Vorticity is a measure of the local rotation of a fluid.
- 17. W. M. Gray, Meteorol. Atmos. Phys. 67, 37 (1998).
- R. M. Zehr, NOAA Technical Report No. NESDIS 61 (NOAA, Washington, DC, 1992).
- 19. E. D. Maloney and D. L. Hartmann, unpublished data.
- 20. D. L. Hartmann and E. D. Maloney, unpublished data. 21. J. Molinari, D. Knight, M. Dickinson, D. Vollaro, S.
- Skubis, Mon. Weather Rev. 125, 2699 (1997).
- 22. R. N. Ferreira and W. H. Schubert, J. Atmos. Sci. 54, 261 (1997).
- 23. The MJO index (the first PC) was normalized to unit variance for use in this figure. Positive values of the index correspond to westerly wind periods. The MJO index has zero mean and includes 1035 positive pentads and 1119 negative pentads. The number of genesis events during westerly periods of amplitude

greater than one is significantly higher than the number during easterly periods of the same amplitude at the 95% confidence level.

- W. M. Gray, *Mon. Weather Rev.* **112**, 1649 (1984).
 C. W. Landsea, G. M. Gray, P. W. Mielke Jr., K. J. Berry, *Weather* **49**, 273 (1994).
- M. C. Bove, J. J. O'Brien, J. B. Eisner, C. W. Landsea, X. Nie, Bull. Am. Meteorol. Soc. 79, 2477 (1998).
- D. E. Waliser, C. Jones, J.-K. E. Schemm, N. E. Graham, J. Climate 12, 1918 (1999).
 - Contribution of Increasing CO₂ and Climate to Carbon Storage by Ecosystems in the United States

30.

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The effects of increasing carbon dioxide (CO_2) and climate on net carbon storage in terrestrial ecosystems of the conterminous United States for the period 1895–1993 were modeled with new, detailed historical climate information. For the period 1980–1993, results from an ensemble of three models agree within 25%, simulating a land carbon sink from CO_2 and climate effects of 0.08 gigaton of carbon per year. The best estimates of the total sink from inventory data are about three times larger, suggesting that processes such as regrowth on abandoned agricultural land or in forests harvested before 1980 have effects as large as or larger than the direct effects of CO_2 and climate. The modeled sink varies by about 100% from year to year as a result of climate variability.

Recent analyses of the global carbon cycle suggest a significant role for terrestrial uptake of CO_2 in the overall budget (1-4). Analyses of atmospheric CO_2 have persistently suggested that this terrestrial uptake is largest in the Northern Hemisphere (2, 3), and one atmospheric analysis suggests that the United States may play a disproportionate role (2). Currently, a number of phenomena contribute to enhanced carbon uptake by ecosystems, including CO_2 fertilization of photosynthesis,

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*To whom correspondence should be addressed. Email: dschimel@bgc-jena.mpg.de and htian@mbl.edu climate, nitrogen deposition, recovery from historical land use, and erosion/sedimentation (4-6). Although preliminary attempts have been made to partition the terrestrial sink among these processes globally, this quantification is currently extremely crude. It is essential to understand the mechanisms controlling carbon exchange today as a basis for prediction and management interventions (7).

28. W. M. Gray, in Meteorology over the Tropical Oceans,

29. E. D. Maloney and D. L. Hartmann, J. Climate 22,

3 December 1999; accepted 2 February 2000

nell, UK, 1979), pp. 155-218.

2387 (1998).

ATM-9873691.

D. B. Shaw, Ed. (Royal Meteorological Society, Brack-

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Here we present results from the Vegetation/Ecosystem Modeling and Analysis Project (VEMAP) aimed at understanding the contribution of ecosystem physiological mechanisms to terrestrial sinks in the conterminous United States during the period 1980-1993. Specifically, we consider how changes in climate and CO₂ concentration affect ecosystem physiology. Three ecosystem models [Biome-BioGeochemical Cycles (Biome-BGC), Century, and the Terrestrial Ecosystem Model (TEM)] that dynamically calculate net carbon storage at a 0.5° imes 0.5° resolution (8) were used. All three models simulated changes to soil and vegetation carbon in natural ecosystems. Century simulated both natural and simulated agricultural ecosystems. To compute complete regional carbon budgets, we identified those grid cells dominated by agriculture (about 40% of the United States) and used the Century agricultural results to simulate net carbon storage in these grid cells for all models. We analyzed the period 1980–1993 to compare our results with spatial atmospheric inverse calculations (2, 3) and inventory-based estimates of forest carbon storage (4, 9, 10). The period spans a range of climatic conditions and includes three El Niño events and the global cooling that followed the eruption of Mount Pinatubo (June 1991).

We used climate information for 1895-1993. Climate information was derived from the National Oceanic and Atmospheric Administration (NOAA) Historical Climate Network (HCN) database (11). Monthly precipitation and mean minimum and maximum temperature information was derived from the NOAA HCN database and other primary, cooperative, and snowpack telemetry (SNOTEL) station data sets (11). Because few station records spanned the entire period, we estimated the spatial autocorrelation structure around each station for its period of record and created a continuous record for 1895-1993 for all stations by geostatistically computing missing anomalies from nearby stations (12). The station data were gridded at 0.5° (latitude by longitude) with a terrainfollowing algorithm (12). Forest types and soils data were as described in VEMAP 1995 (7) with forest distributions remapped with a database derived from satellite observations (13). Agricultural regions were determined from a 1990 land cover inventory of the United States (13) and land management practices (fertilization, planting and harvest dates, irrigation, tillage intensity) from U.S. Department of Agriculture National Resource Inventory data (14, 15). Atmospheric CO₂ data were from Enting et al. (16).

Our ensemble of means for the U.S. $CO_2/$ climate sink in ecosystems for the period 1980–1993 is 0.08 Pg of carbon per year (Table 1). The three models agree within 25% in estimating the continental mean. Comparison of model experiments with observed ver-

Table 1. Net carbon storage in the terrestrial ecosystems of the United States as estimated by three biogeochemical models (petagram of carbon per year).

	Biome-BGC	Century	TEM
1980–1993* Mean	0.081	0.068	0.086
CV† (%) 1988–1992‡	108	108	157
Mean CV† (%)	0.050 249	0.047 172	0.046 293

*The period 1980–1993 is consistent with the inverse analysis of Rayner *et al.* (3). \pm Coefficient of variation. \pm The period 1988–1992 is consistent with the inverse analysis of Fan *et al.* (2).

sus constant CO_2 shows that the bulk of the increase is due to CO_2 fertilization, with the rate of uptake varying with and modulated by climate. Annual net carbon storage per unit area is relatively evenly distributed over the conterminous United States, ranging from 100 kg ha⁻¹ in the Great Plains and the Northeast to 150 kg ha⁻¹ in the Southeast

(Fig. 1). As expected, intermodel variability is higher at the regional level than in the continental total, but the model results remain similar within a factor of 3 and are comparable to inventory-based estimates (9, 10). Agriculture plays a negligible role in modeled current carbon storage, but Century simulations suggest that, with best management



Fig. 1. Net carbon storage for different bioclimatic regions of the United States estimated with all three models for natural ecosystems, with Century results applied in agricultural cells. Histograms show specific model results and the mean (95% confidence interval).



Fig. 2. Interannual variability of continental net carbon storage during 1980–1993 as estimated by three biogeochemical models. Circle and error bar show the mean of three models and standard deviation (\pm) across three models. Dotted line shows mean of three models for the period 1980–1993.

practices, U.S. agriculture can remain a modest sink for decades to come.

The modeled annual carbon sink is highly variable from year to year, ranging from <0.1 Pg of carbon net efflux to >0.2 Pg of carbon uptake, or more than 100% (Fig. 2). The modeled carbon sink for the period 1988–1992 shows much larger year-to-year variations than the entire period (Table 1). The large modeled variation in continental carbon storage is consistent with site-specific flux measurements and atmospheric studies (17), all of which suggest that terrestrial sinks differ dramatically from one year to the next.

This high variability of net carbon storage has several practical implications. First, atmospheric inverse analyses gain in statistical power when they use multiyear averages. In the future, such analyses must be done recognizing that both the magnitude and the subregional spatial distribution of fluxes will vary with the spatiotemporal distribution of climate anomalies. This is especially true if, as is likely, temperature-precipitation anomalies covary with processes (for example, convection) that affect atmospheric transport of CO_2 . Second, any use of annual net carbon storage estimates from eddy covariance or other atmospheric techniques (18) in policy instruments to mitigate climate change or for land management must be cognizant of the volatile nature of this measure. This will certainly affect the notion of commitment periods (18). Third, our results highlight the need for sustained ecological research. Any 2- to 3-year sample within this period could have given a misleading picture of the decadal trend and might have even shown a different sign from the decadal mean. Long-term observations are not merely monitoring, but are needed to recognize fundamental ecological phenomena

The region we simulated is not geographically identical to the regions defined in atmospheric inverse models (2, 3). For perspective, in global simulations with VEMAP models, net carbon storage in the conterminous United States is typically about 60% of the total we model for the region equivalent to the North American domain of Fan et al. (2). Also, the atmospheric signal results from the outcome of all processes, including processes we do not model such as forest regrowth, erosion, and nitrogen deposition. Therefore, we expect, a priori, the atmosphere to show a somewhat larger sink than we model (Table 1). Our estimate is close in magnitude to inventory-based estimates (9, 10) and to some atmospheric estimates (3).

Despite high uncertainty, the inventory estimates tend to be larger than the VEMAP estimate for a CO_2 /climate sink of 0.08 Pg of carbon per year. For example, Brown and

Schroeder (10) estimated 0.17 Pg of carbon per year for eastern U.S. forests (compared with our value of 0.04) (Fig. 1). Birdsey and Heath (9) estimated a U.S. sink of 0.3 Pg of carbon per year, whereas Houghton et al. (4) estimate a range of 0.15 to 0.35 Pg of carbon per year. The effects of intensive forest management and agricultural abandonment on carbon uptake in the United States are probably as large as or larger than the effects of climate and CO_2 . If the total sink is about 0.3 Pg of carbon per year, and the CO₂/climate sink is about 0.1 Pg of carbon per year, other processes such as regrowth on abandoned agricultural and harvested forest lands must cause a sink of about 0.2 Pg of carbon per vear.

This is a different perspective from that given in many global analyses (1, 19). A large role for land use effects is consistent with suggestions from the ecological community in the wake of the Kyoto Protocol (18). The relative roles of physiological (climate, CO_2) changes compared with the direct effects of human domination of ecosystems need to be reassessed as a basis for understanding how the carbon cycle will change in the future.

Despite the discrepancies, the estimates from the VEMAP models are an order of magnitude less than the high atmospherically based estimates of Fan et al. (2). Inventory data also suggest a sink of the order of 0.3 Pg of carbon per year. Thus, the best current information suggests that CO₂ and land use contribute a few tenths of a petagram of carbon uptake each year in the United States. The other hypothesized processes for ecosystem carbon storage (nitrogen deposition and sedimentation) are thought to be of a similar magnitude or smaller in this region (5). Inventory and model results are in conflict with high estimates from atmospheric inverse estimates. The next steps in the quantification of the North American carbon sink will require additional observations (20).

References and Notes

- 1. D. S. Schimel, Glob. Change Biol. 1, 77 (1995).
- 2. S. M. Fan et al., Science 282, 442 (1998).
- P. J., Rayner, I. G. Enting, R. J. Francey, R. Langenfelds, *Tellus* 51B, 213 (1999).
- R. A. Houghton, J. L. Hackler, K. T. Lawrence, *Science* 285, 574 (1999).
- A. Dai and I. Fung, Glob. Biogeochem. Cyc. 7, 599 (1993); E. Holland et al., J. Geophys. Res. 102(D13), 15849 (1997); R. F. Stallard, Glob. Biogeochem. Cyc. 12(2), 231 (1998).
- H. Tian, J. M. Melillo, D. W. Kicklighter, A. D. McGuire, J. Helfrich, *Tellus* 51B, 414 (1999).
- VEMAP Members, *Glob. Biogeochem. Cyc.* 9(4), 407 (1995); T. G. F. Kittel *et al.*, *J. Biogeogr.* 22, 857 (1995).
- We used three biogeochemical models (Biome-BGC, Century, and TEM) for our analyses here. Detailed description of the three models can be found in Running and Hunt [S. W. Running and E. R. Hunt Jr., in Scaling Processes Between Leaf and Landscape Levels, J. R. Ehleringer and C. Field, Eds. (Academic Press,

Orlando, FL, 1993), pp. 141–158] for Biome-BGC, in Parton et al. [W. J. Parton, D. S. Schimel, D. S. Ojima, C. V. Cole, in *Quantitative Modeling of Soil Forming Processes*, R. B. Bryant and R. W. Arnold, Eds. (Special Publication 39, Soil Science Society of America, Madison, WI, 1994), pp. 147–167] for Century; and in (6) for the TEM.

- R. A. Birdsey and L. S. Heath, in Productivity of America's Forest and Climate Change, General Technical Report RM-GTR-271, L. A. Joyce, Ed. (U.S. Department of Agriculture, Forest Service, Rocky Mountain Forest and Range Experiment Station, Fort Collins, CO, 1995), pp. 56–70.
- 10. S. L. Brown and P. E. Schroeder, *Ecol. Appl.* **9(3)**, 968 (1999).
- T. C. Peterson and R. S. Vose, *Bull. Am. Meteorol. Soc.* 78(12), 2837 (1997); T. R. Karl and R. W. Knight, *Bull. Am. Meteorol. Soc.* 79, 231 (1998).
- T. G. F. Kittel et al., Proc. 10th Conference on Applied Climatology, Boston (American Meteorological Society, New York, 1997), pp. 222–229; C. Daly, R. Neilson, D. Philips, J. Appl. Meteorol. 33, 140 (1994).
- T. R. Loveland, J. W. Merchant, D. O. Ohlen, J. F. Brown, *Photogramm. Eng. Remote Sens.* 57(11), 1453 (1991).
- 14. The VEMAP data layer for current vegetation cover was developed from the 1990 1-km EROS Data Center (EDC) Conterminous U.S. Seasonal Land Cover Classification (13) (http://edcwww.cr.usgs.gov/glis/ hyper/guide/landchar). The 154 classes from the 1-km EDC classification were assigned to 33 VEMAP current vegetation types (22 unmanaged classes, 11 managed classes) by VEMAP participants. The reclassed 1-km data were then aggregated to the standard VEMAP 0.5° grid cells by majority.
- 15. Managed vegetation classes were subdivided by regional management practices and growing season length. This was done based on data from the 1995 Cropping Practices Survey (http://usda.mannlib.cornell.edu/data-sets/inputs/93018/) conducted by the Resource Information System Section of the Resources and Technology Division of the Economic Research Service for all crops except sorghum and hay (USDA, 1991 (r 1995i* Cropping Practices Survey, unofficial USDA data files). Sorghum information is from the 1991 Cropping Practices Survey (http://usda. mannlib.cornell.edu/data-sets/inputs/93018/), and hay practices are based on yield data from state-level reports (1995 USDA Annual Crop Summary by state, National Agricultural Statistics Service) (http://usda. mannlib.cornell.edu:70/0/data-sets/crops/9X180/ 97180/2/fieldcrp.txt). Where management practices (irrigation, fertilization, planting or harvest date, tillage practices, or rotation) differed considerably within a land cover class, the land cover class was spatially subdivided to capture these significant differences.
- I. Enting, T. Wigley, M. Heimann, CSIRO Div. Atmos. Res. Technical Pap. 31, 1 (1994).
- M. Goulden et al., Science 271, 1576 (1996); B. H. Braswell, D. S. Schimel, E. Linder, B. Moore III, Science 278, 870 (1997); P. Ciais, P. P. Tans, M. Trolier, J. W. C. White, R. J. Francey, Science, 269, 1098 (1995).
- IGBP Terrestrial Carbon Working Group, Science 280, 1393 (1998).
- 19. M. Cao and I. F. Woodward, Nature 393, 249 (1998).
- J. L. Samiento and S. Wofsy, Eds. A U.S. Carbon Cycle Science Plan (University Corporation for Atmospheric Research, Boulder, CO, 1999).
- 21. VEMAP is sponsored by the Electric Power Research Institute, NASA, and the U.S. Forest Service. Additional support for VEMAP data set development comes from NOAA, NSF, and the National Center for Atmospheric Research (NCAR). NCAR is sponsored by NSF. We thank NOAA's National Climate Data Center for their assistance with climate station data and USDA Natural Resources Conservation Service for access to snowpack data. We thank C. Daly and W. Gibson (Oregon Climate Service, Oregon State University), A. Royle (currently with the U.S. Fish and Wildlife Service), H. Fisher (NCAR), and NSF's Geophysical Statistics Project at NCAR for their crucial role in data set preparation.

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