gas is released to the atmosphere; cooling by 9.1°C (2–4); gas solubility increases for CO_2 by approximately 3.75%. These effects are of opposite sign, and of apparently almost equal magnitude so that the net flux of CO₂ is quite small. For O_2 the above effects are additive and result in a large net flux: exposure of water with low partial pressures of O₂ results in atmospheric O₂ invasion, and cooling further enhances the solubility. These data confirm that both biological and physical cycles of the ocean are important in planetary CO₂ balance. We predict that surface evasion of CO₂ from the North Atlantic will occur in winter. Although a larger net residual southward transport cannot be absolutely ruled out, we believe that the picture we pain? of opposing effects must intuitively be the case and will yield a small net signal.

The calculation of the alkalinity balance shows transports of 71.7×10^6 eq/s northward, and 71.2×10^6 eq/s southward. The difference of 0.5×10^6 eq/s northward is again indistinguishable from zero but suggests that the North Atlantic is a small alkalinity sink. The balance however reflects the interaction of processes quite different than for CO₂; transfer at the air-sea interface does not apply for alkalinity but depends on processes involving CaCO₃ uptake and dissolution and changes in N metabolism (15).

Our calculation, most emphatically, does not mean that oceanic uptake of the fossil fuel signal is small. Transport of fossil fuel CO_2 is taking place in the surface flows and is extractable from the CO_2 flux signal (16). Deep waters in the North Atlantic that have radiochemical and fossil fuel burdens (17) have yet to reach 25°N in other than the deep western boundary current, and thus the present-day balance is artificially poised in time. Interconversion of CO₂ between gaseous and dissolved organic C (18) also occurs, and the magnitude of this cycle is currently controversial (19).

Our estimates of CO₂ transport for a single ocean basin are consistent with the global exchanges between sea and air provided by Pearman et al. (20). These exchanges are calculated to have changed by a factor of 2 in the last 40 years. Unraveling signals such as these is essential for knowledge of the planetary C cycle and will be a principal focus of the Joint Global Ocean Flux Study (JGOFS) (21) in the decade ahead. The large absolute fluxes and the small net signal present enormous challenges to scientists in this field.

(1980). Bryden and Hall made use of International Geophysical Year data (1957) in computing geostrophic flows. Their analysis yielded a heat flux of 1.1 \times 10¹⁵ W. Roemmich and Wunsch (4) in an inverse analysis of a detailed eddy resolving section (1981) obtained a flux of 1.2 × 10¹⁵ W.
3. M. M. Hall and H. L. Bryden, *Deep Sea Res.* 29,

- 339 (1982).
- D. Roemmich and C. Wunsch, ibid. 32, 619 (1985). The oceanic transport of material is vast by continental geochemical standards. The flow of the world's rivers combined is approximately 1.6 Sv; the 80-km-wide Florida Straits alone transports 30 Sv $(1 \text{ Sv} = 10^6 \text{ m}^3/\text{s}).$
- 6. A. Oort and T. von der Haar, J. Phys. Ocean. 6, 781 (1976).
- W. E. Esaias, G. C. Feldman, C. R. McClain, J. W. Elrod, *Eos* 67, 835 (1986).
- 8. T. Platt and S. Sathyendranath, Science 241, 11613 (1988).
- 9. Transient Tracers in the Ocean (TTO): North Atlantic Study (SIO Ref. 86-15, Physical Chemical Oceanographic Data Facility, Scripps Institution of Oceanography, La Jolla, CA, 1981). A. L. Bradshaw and P. G. Brewer, *Mar. Chem.* 23,
- 10. 69 (1988); *ibid.* 24, 155 (1988).
 11. K. M. Johnson *et al.*, *ibid.* 16, 61 (1985).
- 12. S. Rintoul, thesis, Massachusetts Institute of Technology-Woods Hole Oceanographic Institution (1988).
- W. S. Broccker and T. H. Peng, in *Tracers in the Sea* (Eldigio Press, Palisades, NY, 1982), pp. 1–690.
 A. F. Bunker, *Mon. Weather Rev.* 104, 1122 (1976).

- P. G. Brewer and J. C. Goldman, *Limnol. Oceanogr.* 21, 108 (1976); J. C. Goldman and P. G. Brewer, *ibid.* 25, 352 (1980).
- P. G. Brewer, Geophys. Res. Lett. 5, 997 (1978).
 See papers in the TTO collection, J. Geophys. Res.
- 90, 6903 (1985) and in (9) 18.
- Y. Sugimura and Y. Suzuki, Mar. Chem. 24, 105 (1988); G. Jackson, Oceanography 1, 28 (1988). E. T. Peltzer and P. G. Brewer, unpublished data.
- DOC data obtained show a high surface value of 175 µmol of C per kilogram and a low of 75 µmol of C per kilogram. The data are linearly related with the apparent O₂ utilization signal. 20. G. L. Pearman, P. Hyson, P. J. Fraser, J. Geophys.
- Res. 88, 3581 (1983).
- 21. The Joint Global Ocean Flux Study (JGOFS) is a decade-long international experiment in which satellite observations of ocean color and sampling in oceanic cruises are coordinated.
- 22. We thank D. Shafer, A. Fleer, N. Hayward, and A. Bradshaw for help, and C. Wunsch and H. Bryden for critical comment. The thesis of S. Rintoul and his analysis of the nutrient cycles further influenced our work. We are grateful to A. Edwards for manuscript preparation. Our work was supported by the National Science Foundation under grants OCE87-1461 (PGB); by a Bourse Lavoisier from the Ministère des Affaires Etrangères (C.G.); and by a grant from the WHOI Coastal Research Center to support the visit of D.D. Contribution 7109 from the Woods Hole Oceanographic Institution.

24 May 1989; accepted 1 September 1989

A Devonian Spinneret: Early Evidence of Spiders and Silk Use

WILLIAM A. SHEAR, JACQUELINE M. PALMER, Jonathan A. Coddington, Patricia M. Bonamo

A nearly complete spider spinneret was found in Middle Devonian rocks (about 385 to 380 million years old) near Gilboa, New York. This is the earliest evidence yet discovered for silk production from opisthosomal spigots, and therefore for spiders. Two previously known Devonian fossils described as spiders lack any apomorphies of the order Araneae and are probably not spiders. The spigots of the Devonian spinneret resemble those of members of the living suborder Mesothelae, but the number of spigots and their distribution are like those of members of the suborder Opisthothelae, infraorder Mygalomorphae. The Devonian spider belonged to a clade that may be the sister group of all other spiders, of Mesothelae, or of Opisthothelae.

PIDERS (ARTHROPODA: CHELICERata: Araneae) are among the most important terrestrial predatory animals. Among the arachnids, they alone produce silk from opisthosomal (abdominal) glands that open through modified setae called spigots, which in turn are located on reduced abdominal appendages, the spinnerets. This character complex is the most diagnostic apomorphy of spiders. We report here on the earliest evidence yet discovered in the fossil record of spinnerets, of spiders themselves, and of silk production by animals.

Although two spider fossils have been reported from the Devonian Period, in neither of these cases can any apomorphies of the order Araneae be demonstrated. Paleocteniza crassipes (1), from the Lower Devonian (404 million years old?) Rhynie Chert, is a minute, crumpled exoskeleton that is undoubtedly arachnid, but is more likely from one of the trigonotarbids that are the most abundant animals in that deposit. Spinnerets, characteristic patterns of leg jointing, eye arrangement, and other spider apomorphies that are potentially present even in very small, immature animals cannot be detected in this fossil or are certainly not there (2). Archaeometa devonica (3), from the slightly later Alken-an-der-Mosel, West Ger-

REFERENCES AND NOTES

^{1.} P. G. Brewer and D. Dyrssen, paper presented at the First Joint Global Ocean Flux Study (JGOFS) meeting, Paris, 17 to 20 February 1987.

^{2.} H. L. Bryden and M. M. Hall, Science 207, 884

W. A. Shear, Department of Biology, Hampden-Sydney College, Hampden-Sydney, VA 23943, and American Museum of Natural History, New York, NY 10024. J. M. Palmer, Museum of Comparative Zoology, Har-vard University, Cambridge, MA 02138.

J. A. Coddington, Department of Entomology, National Museum of Natural History, Smithsonian Institution, Washington, DC 20560. P. M. Bonamo, Center for Evolution and the Paleoenvi-ronment, State University of New York, Binghamton,

NY 13901.



Fig. 1. Fossil spinneret, slide 334-1b-AR34. The greatest length of the specimen (not including the terminal spigots) is 0.94 mm. Note the numerous spigots scattered along the anatomically median surface. Magnification, $\times 40$.

many, judging from the published photographs, is not a spider and perhaps not even an animal fossil, rather a vertebrate coprolite. Again, no spider apomorphies are visible on the specimen, which is simply an elongate blob with vague cross-striations at one end. In our opinion, it must be rejected as a possible spider fossil.

The earliest known terrestrial arachnids in North America occur in Middle Devonian rocks near Gilboa, New York (4, 5). An extraordinarily diverse fauna of arthropods, some not yet identified, are found in this *Konservat Lagerstätte* (a fossil deposit remarkable for fine preservation, preservation of an entire community, or both). Among recently obtained specimens is a single spider spinneret.

Recent views of spider evolution (6) divide the order Araneae into two suborders. Mesothelae includes a small number of species today restricted to southeast Asia, Indonesia, and Japan; they are united by a number of synapomorphies, including a peculiar sense organ between the tibiae and metatarsi of the legs (7). Mesotheles are better known for their primitive characters, including an externally segmented opisthosoma and the possession of eight (rarely seven) spinnerets, which are located not at the end of the opisthosoma, but near the middle of its ventral surface. Suborder Opisthothelae includes all other spiders, in which the number of spinnerets has been reduced to six, four, or two and moved to the posterior end of the opisthosoma, which is not externally segmented. Within this group, Mygalomorphae ("tarantulas" in the North American sense) have lost all vestiges of the anterior median spinnerets, while Araneomorphae carry a cribellum (repeatedly lost in many lines) homologous to the anterior median spinnerets of mesotheles, and have chelicerae rotated to the labidognath position, so that the fangs point toward one another.

Aside from the Devonian examples mentioned above, all Paleozoic fossil spiders





Fig. 2. Distal portion of the fossil spinneret, showing denser clustering of spigots at tip. Magnification, $\times 200$.



Fig. 3. Posterior median spinnerets of Liphistius malayanus (Mesothelae: Liphistiidae). Note the single, terminal spigots, and the scaly cuticle. Scanning electron micrograph (SEM), ×92.

come from Pennsylvanian rocks, and with two possible exceptions (8) are mesotheles, though spinnerets are not preserved in the majority and this assignment by paleontologists has been based on the combination of spider-like general morphology and a segmented opisthosoma. Some of these fossils are not spiders (9). A gigantic Carboniferous arachnid from Argentina has been assigned to Araneae (10) but may represent an unnamed order or a ricinuleid (11). In any case, this fossil, *Megarachne servinei*, suggests undetected Paleozoic araneid or arachnid diversity.

The Devonian spinneret (Figs. 1 and 2) is nearly complete, consists of a single article, and carries 19 or 20 spigots that are in most ways characteristic of mesotheles. The spigots are arrayed along the medial surface and are more densely clustered distally. Slit sense organs and setal sockets are scattered over the cuticle between spigots, and a few of the sockets retain setae, which may be either smooth or serrate. The cuticle itself has a scaly appearance, as does that of living mesotheles. However, in mesotheles the large lateral spinnerets of each pair are pseudosegmented, with spigots in ranks of two, three, or four on the mesal surface of a pseudosegmental ring, and the smaller, single-articled median ones bear only a single spigot (Liphistius maylayanus, Fig. 3). Because the Devonian spinneret is not pseudosegmented,



Fig. 4. Posterior median spinneret of *Neocteniza* sp. (Opisthothelae: Mygalomorphae: Idiopidae). The spigots are numerous, occupy the median surface, and are more densely clustered near the tip. SEM, \times 92.

yet bears more than one spigot, it could not have come from a mesothele spider similar to those living today.

We have ruled out araneomorph spiders since the spigots of their spinnerets are strongly differentiated from one another and from those of mygalomorph spiders in characteristic ways (12), and all spigots on the fossil specimen are of the same size and shape.

Mygalomorph spiders have single-articled posterior median spinnerets with numerous spigots (Neocteniza sp., Fig. 4) arranged as in the fossil. The presence of undifferentiated, or only slightly differentiated, spigots that are more densely clustered near the tip of the spinneret is consistent with mygalomorph spider posterior median spinneret anatomy. However, both mygalomorph and araneomorph (but not mesothele) spinnerets have peculiar nipple-shaped structures called tartipores, which represent the positions of spigots in previous instars (12). Tartipores are not present on the Devonian spinneret. In addition, mygalomorph spinnerets usually have two types of spigots present. The form of the spigots themselves does not, in detail, agree with that of mygalomorph spigots (Table 1).

Mesothele spigots (Fig. 5) are uniform in morphology, with a broad, conical base and a long, gradually tapering, unsculptured distal shaft that merges smoothly into the base. The spigots of our fossil (Fig. 6) are of this type.

Mygalomorph spigots usually have an articulated shaft, which joins the base by means of a well-defined, sleevelike fold. At least the distal third of the shaft is sculptured. However, the rastelloid clade of mygalomorphs have nonarticulated shafts and extremely fine sculpture, visible only when viewed with a scanning electron microscope. Diagenetic changes in the fossil spinneret may have made it impossible to resolve such fine detail as the distal shaft sculpture of *Neocteniza*. Considering the absence of tarti
 Table 1. Comparison of spinnerets.

Character	<i>Liphistius</i> (mesothele) posterior median spinneret	Devonian fossil spinneret	Mygalomorph posterior median spinneret
Spigot arrangement	Single apical spigot	19–20 on mesal side of spinneret, not ranked, clustered at tip	Numerous on mesal side of spinneret, not ranked, clustered at tip
Spigot types	One	One	Rarely one, usually two
Cuticle texture	Scaly	Less pronounced scales	Slightly scaly
Shaft sculpture	Absent	Apparently absent	Present on at least distal third
Shaft-base union Tartipores	Smoothly graded Absent	Smoothly graded Absent	Collar-like articulation Present

pores and the possibility that distal sculpture is absent, not eroded by postmortem changes, the spigots are more like mesothele spigots than mygalomorph ones.

As already discussed above, the combinations of apomorphies found in spinnerets of the three living clades would seem to exclude the fossil from all of them. The question then becomes placement of the Devonian spider as a sister group of one, two, or all of these clades. The presently accepted three-taxon statement for the groups of spiders so far discussed is (Mesothelae (Mygalomorphae (Araneomorphae))). The spider that bore the fossil spinneret is probably not a member of the sister group of either Araneomorphae or Mygalomorphae, because to place it in either of those positions would require the ad hoc secondary loss of tartipores in the fossil clade. Thus the fossil may be a representative of the sister group to all other spiders, to Mesothelae, or to Opisthothelae. Additional evidence from other parts of the Gilboa spider is required to further refine its position, since all observable character states of the spinneret are plesiomorphic.

However, the early appearance of every physical modification required to produce silk at a level of sophistication paralleling that of some modern spiders is striking. If relatively constant rates of evolution are assumed, it suggests a long period of pre-Middle Devonian evolution for spiders and their relatives, and that even Devonian fossils will not shed much light on the origins of spider spinning (13). While mesotheles may have achieved their modern form by the Pennsylvanian, definitively opisthothele fossils do not appear until the Mesozoic [Triassic (14), Jurassic (15), Cretaceous (16)], and nearly all of these fossils can be assigned to families still extant-they are in every detail (the preservation is exquisite) the equivalent of living species.

To what use Devonian spiders put their silk is unclear. Living mesotheles do not



Fig. 5. Terminal spigot on posterior median spinneret of L. malayanus. SEM, ×510.



Fig. 6. Spigot base from fossil spinneret. Oil immersion, N Nomarski interference contrast optics, ×1000.

make aerial webs but use silk only as a burrow-and-door lining, as trip lines extending from the mouth of the burrow, and as the material for egg sacs. However, when spiders make trap doors, there are specific adaptations present to shorten and broaden the spinnerets (17), which are not present in our specimen. Rudimentary aerial webs are made by a few mygalomorphs (18), and many araneomorphs weave highly derived ones (19), a habit that may be correlated with their well-differentiated spigots. Flying insects, against which aerial webs would have been directed, do not appear in the fossil record until much later [Carboniferous: Namurian (20)] but may have had a long history previous to that appearance.

The report (21) of an archaeognath insect from the Lower Devonian (Emsian), and the presence in the later (Givetian) Gilboa fauna of similar material (4) establishes an early origin for insects. Devonian material of winged (pterygote) insects may well be found in the near future.

The Devonian spider, therefore, was perhaps a sit-and-wait, tunnel or tube-dwelling predator on cursorial arthropods, but may just possibly have made an aerial web.

REFERENCES AND NOTES

- 1. S. Hirst, Annu. Mag. Nat. Hist. 9, 455 (1922). 2. W. A. Shear, unpublished observations. The speci-
- men was studied intensively using Nomarski optics; 35 serial optical sections were photographed, digitized, and reconstructed on a computer. The result-ing image could be examined from all angles; no spider apomorphies were found.
- 3. L. Størmer, Senckenb. Lethaea 57, 121 (1976).
- W. A. Shear et al., Science 224, 492 (1984).
 W. Shear, P. Selden, W. D. I. Rolfe, P. Bonamo, J. Grierson, Am. Mus. Novit., no. 2901, 1 (1987); R. Norton, P. Bonamo, J. Grierson, W. Shear, J. Paleontol. 62, 259 (1988); W. Shear and P. Bonamo, Am. Mus. Novit., no. 2927, 1 (1988). N. Platnick and W. Gertsch, Am. Mus. Novit., no.
- 6. 2607, 1 (1976); R. Raven, Bull. Am. Mus. Nat. Hist. 182, 1 (1985)
- N. Platnick and P. Goloboff, J. NY Entomol. Soc. 7. 93, 1265 (1985)
- 8. W. A. Shear, unpublished observations. Archaeometa nephilina Pocock and Arachnometa tuberculata Petrunkevitch [both Late Carboniferous; British Museum of Natural History (BMNH), specimens In15863, In31259, and In13914, respectively] may be araneomorphs.
- W. A. Shear, unpublished observations. Procteniza brittanica Petrunkevitch (Late Carboniferous; BMNH specimen In22834), for example, has the ventral opisthosomal surface very well preserved, but shows no spinnerets. It is probably an amblypy-
- gid. 10. M. Hünicken, Bol. Acad. Nacl. Ciencias, Cordoba, Argentina 53, 317 (1980).
- 11. P. Selden, personal communication.
- J. Coddington, J. Arachnol. 17, 75 (1989); J. Ko-voor, in Ecophysiology of Spiders, W. Nentwig, Ed. (Springer-Verlag, New York, 1986), pp. 160–186. Schultz, Biol. Rev. 62, 89 (1987)
- 14. J.-C. Gall, Mem. Service Carte Geol. d'Alsace Lorraine **34**, 1 (1971).
- K. Eskov, N. Jahrb. Geol. Palaeontol. Monatsh. 11, 645 (1984); N. Jahrb. Geol. Palaeontol. Abh. 175, 81 1987)
- 16. P. Selden, personal communication; J. McAlpine and J. Martin, Can. Entomol. 101, 827 (1969).
- 17. R. Raven, Bull. Am. Mus. Nat. Hist. 182, 12 (1985). 18. F. Coyle, in Spiders: Webs, Behavior, and Evolution, W. A. Shear, Ed. (Stanford Univ. Press, Stanford, CA, 1986), pp. 274–279, figures 10.6 and 10.7; Bull. Am. Mus. Nat. Hist. 187, 205 (1988).
- 19. W. A. Shear, in Spiders: Webs, Behavior, and Evolution, W. A. Shear, Ed. (Stanford Univ. Press, Stanford, CA, 1986), pp. 364–400; W. Eberhard, *ibid.*, pp. 72–75; J. Coddington, *ibid.*, pp. 326–334.
 20. R. Wootton, Annu. Rev. Entomol. 26, 320 (1981).
- C. C. Labandeira, B. S. Beall, F. M. Hueber, Science 242, 913 (1988)
- 22. We thank A. Robble, who prepared the specimen, N. Platnick, who loaned material of Liphistius, and P. Selden, for commenting on the manuscript, and for discussions with W.A.S. of fossil spiders. The comments of two anonymous reviewers improved the manuscript. Supported by NSF grants BSR 85-084-42 and BSR 88-180-27 to W.A.S. and P.M.B., and by a research grant from the Power Authority of New York State to P.M.B.

30 June 1989; accepted 29 August 1989