

- in (5) and in R. Morlan [*Taphonomy and Archaeology in the Upper Pleistocene of the Northern Yukon Territory: A Glimpse of the Peopling of the New World* (Archaeological Survey of Canada Mercury Series Paper 94, National Museum of Man, Ottawa, 1980)].
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First Radioisotope (Potassium-Argon) Age of Marine Neogene Rionegro Beds in Northeastern Patagonia, Argentina

Abstract. An average age of 9.41 million years was obtained from radioisotope (potassium-argon) age determinations of three glass concentrates of a tuff from the upper part of the marine Rionegrense at Punta Cracker in Golfo Nuevo, Argentina. This age correlates with the Tortonian marine stage of Europe and the Chasicooan Land Mammal Age of South America.

Ameghino (1) proposed the first classification of the marine Neogene of Argentina. He (1) recognized two major subdivisions, the Entrerriense (~ Mio-Pliocene) and the Rionegrense (~ Pliocene) (2). Beds of these ages crop out discontinuously over a broad area of northeastern Argentina from the Paraná basin (61°W, 32°S) south to the vicinity of Península Valdés (Fig. 1). Eustatic sea-level changes and gentle warping of this part of Argentina during the late Cenozoic have resulted in a number of transgressive cycles. The discontinuous nature of the outcrops and the similarity in the lithology of these deposits make correlations difficult. As a result, the age and stratigraphic succession of these beds is disputed (3, 4). Earlier age assignments of these beds were based on either comparisons of the molluscan faunas with coeval European faunas (5) or by the use of Lyellian percentages (6). This has led to age assignments ranging from Oligocene to Pleistocene (1, 3, 4, 6, 7).

Although rocks referable to these stages are widespread throughout northern Chubut and Buenos Aires provinces, the most complete sequence is exposed along the sea cliffs around Península

Valdés. The rocks in this region are nearly flat-lying and consist of nearshore interbedded sandstones and tuffaceous siltstones. The sequence at Punta Cracker on the southwestern shore of Golfo Nuevo (Fig. 1) represents a prograding nearshore-beach facies changing from shallow marine to beach-lagoon deposits. Earlier investigators (4, 5) have referred the shallow marine facies to the

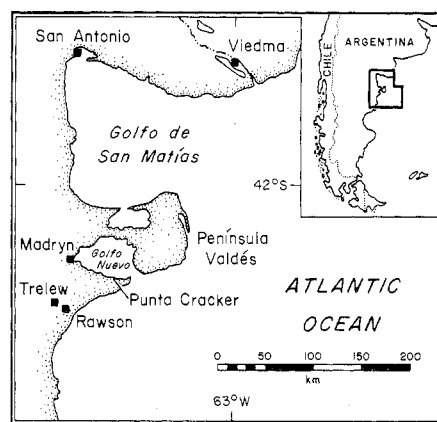


Fig. 1. Map of the northeastern coast of Patagonia, showing the location of Punta Cracker.

Table 1. Analytical data for vitric concentrates of tuff dated at Punta Cracker (8).

Sample number	Sample weight (g)	K (%)	Radiative ^{40}Ar ($\times 10^{-11}$ mole/g)	Atmospheric ^{40}Ar (%)	Age ($\times 10^6$ years)
KA 3509	1.0193	3.48	5.52	54	9.11 \pm 0.1
KA 3633	5.7335	3.23	5.38	46	9.56 \pm 0.3
KA 3510	0.3301	3.43	5.69	73	9.55 \pm 0.3

Entrerriense and the beach-lagoonal facies to the Rionegrense.

Samples of a whitish tuff 2 m thick from near the top of the Rionegrense horizon at Punta Cracker were collected and dated by the ^{40}K - ^{40}Ar method (8). Three glass concentrates gave dates ranging from 9.11 (sample KA 3509) to 9.56 (sample KA 3633) $\times 10^6$ years (mean = 9.41×10^6 years) (Table 1). These dates correlate with the Late Miocene Tortonian marine stage in Europe (9) and the Chasicooan Land Mammal Age in South America (10). Because of the cyclic nature of the Neogene transgressions in northeastern Argentina, it is probable that deposits elsewhere referred to the Rionegrense may be older or younger than those at Península Valdés.

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8. The ^{40}K - ^{40}Ar dates are based on glass concentrates of a water-laid tuff containing abundant ripple marks. The concentrates were composed of primary volcanic glass shards obtained by heavy liquid separation with bromoform. This separation technique effectively removed fine-grained biotite flakes and other detrital material. Each date was obtained on a different sample of tuff, and in each case separate preparations were used. These facts and the consistently similar dates demonstrate repeatability and high reliability for the age of this tuff. Calculations are based on the decay constants $^{40}\text{K}\lambda\beta = 4.962 \times 10^{-10} \text{ year}^{-1}$ and $\lambda^{40}\text{K}e + ^{40}\text{K}e = 0.581 \times 10^{-10} \text{ year}^{-1}$ and on the isotope abundance $^{40}\text{K} = 0.01167$ percent of the total potassium.
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